Total soil organic carbon increases but becomes more labile after afforestation in China’s Loess Plateau

Qingyin Zhang, Xiaoxu Jia, Xiaorong Wei, Mingan Shao, Tongchuan Li, Qiang Yu

State Key Laboratory of Soil Erosion and Dryland Farming on the Loess Plateau, Northwest A&F University, Yangling 712100, China
Key Laboratory of Ecosystem Network Observation and Modeling, Institute of Geographic Sciences and Natural Resources Research, Chinese Academy of Sciences, Beijing 100101, China
College of Resources and Environment, University of Chinese Academy of Sciences, Beijing 100190, China
School of Life Sciences, University of Technology Sydney, P.O. Box 123, Broadway, NSW 2007, Australia

Abstract

Afforestation of cropland is recommended as an effective approach to enhance soil organic carbon (SOC) sequestration and labile organic C fractions. However, the stabilization of SOC and its labile organic C fractions on the Loess Plateau is largely unknown. Our objective was to quantify total SOC concentration and labile organic C fractions in the 0–20 cm soil depth for four land use types on the Loess Plateau, including cropland and three afforested areas (composed of R. pseudoacacia forests, P. tabuliformis forests, and R. pseudoacacia + P. tabuliformis mixed forests). Total SOC concentration, particulate organic C (POC), dissolved organic C (DOC), microbial biomass C (MBC), and potassium permanganate-oxidizable C (KMnO4-C) were measured. Carbon management index (CMI) was also calculated. Afforestation showed a significant positive effect on total SOC and labile organic C fractions, compared with cropland. Afforestation with R. pseudoacacia, P. tabuliformis, and R. pseudoacacia + P. tabuliformis significantly increased SOC by 57.4%, 22.2%, and 44.4% in the 0–5 cm soil layer; and similar increases were observed in the 5–10 cm and 10–20 cm layers. Similar trends to those observed for SOC in response to afforestation were also seen for DOC, MBC, and KMnO4-C. Afforestation with R. pseudoacacia resulted in the highest total SOC concentrations and labile organic C fractions among the three afforestation treatments. These findings suggested that although afforestation can significantly promote total SOC accumulation, especially with R. pseudoacacia, SOC may become more labile following afforestation in the future.

1. Introduction

Forests are among the most productive and ecologically valuable ecosystems in the world (Richter et al., 1999; Foley et al., 2005; Miles and Kapos, 2008). They provide critical ecosystem services, including carbon (C) storage, biodiversity conservation, water purification, and erosion control (Lal, 2005; Metz et al., 2007; Canadell and Raupach, 2008). One of the most important ecological functions of forests is their role in global C cycles (Houghton and Hackler, 1999; Six et al., 2002; Guo and Gifford, 2002). Soil organic C (SOC) is the largest C stock in the terrestrial ecosystem (Batjes, 1996). Soil organic C stocks are controlled by the balance between C inputs and outputs from soils, and afforestation may influence C input as well as output fluxes from ecosystems (Guo and Gifford, 2002; Lal, 2004). Reports from around the world indicate that afforestation can increase soil C sequestration by simultaneously decreasing C loss from decomposition and erosion (Laganière et al., 2010; Xiao et al., 2015; Deng et al., 2016; Li et al., 2018). However, some studies have reported negligible effects of afforestation on C sequestration. (Rytter, 2016; Chen et al., 2017). The positive effects of afforestation are particularly important in arid and semi-arid regions where fragile ecosystems can suffer severe soil degradation and erosion.

Because soil C accumulation plays an important role in climate change mitigation, numerous studies have reported the dynamics of total organic C and labile organic C fractions (e.g. dissolved organic C (DOC), microbial biomass C (MBC), particulate organic C (POC), and potassium permanganate-oxidizable C (KMnO4-C)) with regard to afforestation in different regions (Blair et al., 1995; Wang et al., 2009; Shang et al., 2018; Bargali et al., 2018; Kooch et al., 2019, Pang et al., 2019). Soil total C generally has been reported to increase following afforestation. For example, Martens et al. (2003) found that soil C accumulated at an average rate of 0.62 Mg ha−1 yr−1 during cropland conversion to forest in Central America. Globally, Post and Kwon (2000) found that the average rate of soil C accumulation for forest...
established on cropland was 0.34 Mg ha$^{-1}$ yr$^{-1}$. Guo and Gifford (2002) concluded that soil C stocks significantly increased when cropland was converted to tree plantation (+18%) and secondary forest (+53%). However, Vesterdal et al. (2002) observed that afforestation on cropland did not lead to an increase in soil organic C over a 30-yr period, but led to a redistribution of SOC in the soil profile. Some findings about the changes of soil labile organic C following afforestation have also been reported. For instance, in a temperate forest, Kooch et al. (2019) reported that a tree plantation showed higher POC and DOC than a mixed natural forest. Natural restoration resulted in markedly higher SOC and labile organic C than artificial afforestation in the karst regions of southwestern China (Pang et al., 2019). Bargali et al. (2018) reported that mixed grassland-pine forest was better in sustaining the soil MBC than pure oak or pine forest. They indicated that tree species is a key factor in determining SOC fractions along with vegetation succession. The inconsistencies of soil C dynamics observed in these studies are likely a result of multiple factors, including climate, soil type, soil depth, and tree species (Paul et al., 2002; Lal, 2004; Laganière et al., 2010; Li et al., 2012). Even though several studies have considered the changes of total organic C and labile organic C in pure forest types or different plantations, some recent studies comparing pure vs. mixed stands obtained different results for labile organic C (Shang et al., 2018; Bargali et al., 2018; Kooch et al., 2019; Pang et al., 2019). Moreover, few studies have assessed the possible effects of different tree species mixtures in China’s Loess Plateau.

China implemented the “Grain for Green” programs with the objective of restoring degraded cropland to forest and grassland. *Robinia pseudoacacia* was chosen as one of the tree species because of its N$_2$-fixation capability, an important pathway for atmospheric N$_2$ input into terrestrial ecosystems. Additionally, the species *P. tabuliformis* (non-N$_2$-fixing) was also used as a prominent artificial afforestation forest in the Loess Plateau. Consequently, this project has significantly increased total organic C accumulation in the soil profile and enhanced total organic C storage for over a decade on the Loess Plateau (Zhang et al., 2019). The study site features a temperate, semi-humid climate, with a mean annual temperature of 11.5 °C and a mean annual precipitation of 592 mm, 70% of which occurs during the growing season from June to September. The soil is mainly Gleyic Phaeozems (World Reference Base for Soil Resources), with soil texture of 11% sand, 20% clay, and 69% silt (Zhang et al., 2018). In the study area, there are deciduous broad-leaf and evergreen coniferous forests characterized by *R. pseudoacacia* Linn. and *P. tabuliformis* Carr., which dominate artificial afforestation forests. Maize (*Zea mays* L.) and wheat (*Triticum aestivum* L.) are the predominant crops in this region.

### 2. Materials and methods

#### 2.1. Study area

This study was carried out at the Yehe National Forestry Center of Fufeng County, Shaanxi Province, China (34.55°N, 107.90°E; 1080 m a.s.l.). The study area is part of the Qishui watershed which belongs to the hilly and gully zone of the Loess Plateau (Zhang et al., 2019). The study site features a temperate, semi-humid climate, with a mean annual temperature of 11.5 °C and a mean annual precipitation of 592 mm, 70% of which occurs during the growing season from June to September. The soil is mainly Gleyic Phaeozems (World Reference Base for Soil Resources), with soil texture of 11% sand, 20% clay, and 69% silt (Zhang et al., 2018). In the study area, there are deciduous broad-leaf and evergreen coniferous forests characterized by *R. pseudoacacia* Linn. and *P. tabuliformis* Carr., which dominate artificial afforestation forests. Maize (*Zea mays* L.) and wheat (*Triticum aestivum* L.) are the predominant crops in this region.

#### 2.2. Soil sampling and processing

The three forests in our study were: (i) *R. pseudoacacia* forests, 20 years old; (ii) *P. tabuliformis* forests, 28 years old; and (iii) *R. pseudoacacia* + *P. tabuliformis* mixed forests, 25 years old. A cropland site, harvested once a year, was used as the control because the three afforested areas had been converted from croplands. For the cropland area, both nitrogen and phosphorus (*urea + P$_2$O$_5$, dry form*) were applied at the rates of 100 kg N ha$^{-1}$ yr$^{-1}$ and 50 kg P ha$^{-1}$ yr$^{-1}$ annually at the end of May. Three blocks (distance of 500 m from each other) with similar climatic conditions, soil texture (Table 1), and parent material were randomly selected in the cropland area and each afforested area, which may minimize confounding site factor effects.

### Table 1

<table>
<thead>
<tr>
<th>Land use</th>
<th>Soil layers (cm)</th>
<th>Sand (%)</th>
<th>Silt (%)</th>
<th>Clay (%)</th>
<th>BD (g cm$^{-3}$)</th>
<th>Total organic C concentration (g kg$^{-1}$)</th>
<th>Total organic C stocks (kg m$^{-2}$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Cropland</td>
<td>0–5</td>
<td>10.2(0.3)</td>
<td>68.9(3.5)</td>
<td>20.9(1.2)</td>
<td>1.28(0.1)</td>
<td>11.5(1.1)c</td>
<td>7.36(0.05)b</td>
</tr>
<tr>
<td></td>
<td>5–10</td>
<td>11.6(0.4)</td>
<td>69.6(4.1)</td>
<td>18.8(0.9)</td>
<td>1.36(0.1)</td>
<td>8.4(0.8)</td>
<td>5.71(0.04)b</td>
</tr>
<tr>
<td></td>
<td>10–20</td>
<td>11.8(0.5)</td>
<td>70.5(2.9)</td>
<td>17.7(1.4)</td>
<td>1.38(0.1)</td>
<td>7.6(0.6)</td>
<td>10.49(0.06)b</td>
</tr>
<tr>
<td>RF</td>
<td>0–5</td>
<td>7.6(0.4)</td>
<td>70.9(3.7)</td>
<td>21.5(1.8)</td>
<td>1.18(0.2)</td>
<td>21.8(1.3)a</td>
<td>12.86(0.09a)</td>
</tr>
<tr>
<td></td>
<td>5–10</td>
<td>9.8(0.8)</td>
<td>68.1(2.4)</td>
<td>22.1(2.0)</td>
<td>1.29(0.1)</td>
<td>12.9(1.2)a</td>
<td>8.52(0.07a)</td>
</tr>
<tr>
<td></td>
<td>10–20</td>
<td>12.1(0.9)</td>
<td>69.3(4.1)</td>
<td>18.6(1.2)</td>
<td>1.42(0.1)</td>
<td>9.3(0.7a)</td>
<td>13.77(0.09a)</td>
</tr>
<tr>
<td>PF</td>
<td>0–5</td>
<td>8.8(0.7)</td>
<td>74.6(5.6)</td>
<td>16.4(1.2)</td>
<td>1.23(0.1)</td>
<td>14.5(0.9)b</td>
<td>8.92(0.10b)</td>
</tr>
<tr>
<td></td>
<td>5–10</td>
<td>9.2(0.8)</td>
<td>73.5(4.1)</td>
<td>17.3(1.4)</td>
<td>1.38(0.1)</td>
<td>8.8(0.8b)</td>
<td>5.93(0.07b)</td>
</tr>
<tr>
<td></td>
<td>10–20</td>
<td>10.7(0.8)</td>
<td>71.9(6.5)</td>
<td>17.4(1.3)</td>
<td>1.45(0.1)</td>
<td>8.7(0.4a)</td>
<td>11.17(0.07b)</td>
</tr>
<tr>
<td>RPF</td>
<td>0–5</td>
<td>8.2(0.7)</td>
<td>73.5(5.8)</td>
<td>18.7(1.1)</td>
<td>1.19(0.1)</td>
<td>19.1(1.7a)</td>
<td>11.36(0.08a)</td>
</tr>
<tr>
<td></td>
<td>5–10</td>
<td>9.5(0.6)</td>
<td>72.8(4.9)</td>
<td>17.7(0.9)</td>
<td>1.26(0.1)</td>
<td>12.6(1.1a)</td>
<td>7.94(0.05a)</td>
</tr>
<tr>
<td></td>
<td>10–20</td>
<td>11.8(1.1)</td>
<td>71.2(5.2)</td>
<td>17.0(0.8)</td>
<td>1.42(0.1)</td>
<td>9.2(0.8a)</td>
<td>13.06(0.10a)</td>
</tr>
</tbody>
</table>
Three replicated plots (5 m × 5 m for cropland, 20 m × 20 m for the three afforested areas) were randomly chosen in each block.

Soil samples were taken using a 5.0 cm diameter auger in each plot in June 2018 from five locations (including the four corners and the center). This was the time just before *R. pseudoacacia* trees shed their leaves. Each soil sample was divided into three layers (0–5 cm, 5–10 cm, and 10–20 cm) and the five samples from each layer were composited into one sample. After roots and other plant debris were removed, the samples were divided into two parts. One part of the fresh soil sample was sieved moist (< 2 mm) and kept at 4 °C for DOC and MBC analysis conducted within two weeks of sampling. The other portion of each soil sample was air-dried and kept at room temperature for analysis of SOC concentration, POC, and KMnO₄-C. Three additional soil cores per plot were taken from the three soil layers using stainless steel cylinders (5 cm inner diameter and 5 cm height) for bulk density (BD) analysis. The main physical and chemical properties of the soil for the four land use types are shown in Table 1. In addition, three sampling points were randomly selected in each plot using a 1 m × 1 m quadrat for the aboveground biomass measurements of understory vegetation on the same date. All of the biomass samples were then oven-dried at 70 °C to a constant weight and weighed to determine aboveground biomass.

### 2.3. Soil analysis and data calculation

Soil particle-size distributions were determined using the laser-diffraction method with a Mastersizer 2000 particle-size analyzer (Malvern Instruments, Worcestershire, UK). Total SOC was determined by the dichromate wet oxidation method (Nelson and Sommers, 1982), which may result in smaller values than the dry combustion method. However, this difference will not affect our comparative study among the afforestation treatments. For DOC, the moist soil was extracted with 50 mL of 0.5 mol L⁻¹ K₂SO₄ for 1 h. The extracts were filtered through a 0.45 μm membrane filter and analyzed using a Multi 3100 N/C analyzer (Analytik Jena, Germany) (Jones and Willett, 2006). MBC was measured by the fumigation-extraction method (Vance et al., 1987). In brief, extracted C concentration was determined using a Multi 3100 N/C analyzer (Analytik Jena, Germany). Extracted C concentration was calculated as: MBC (mg g⁻¹) = (fumigated C-non-fumigated C)/0.38 (Vance et al., 1987). PO2 was measured as described by Cambardella and Elliott (1992). In brief, 20 g of air-dried soil and 70 mL of Na hexametaphosphate (5 g L⁻¹) were shaken on a centrifuge for 18 h. The soil suspension was poured over a 53-μm sieve and the retained coarse fraction was rinsed with a weak stream of distilled water. All materials remaining on the sieve were washed into a dry dish, oven-dried at 60 °C for 48 h, and ground to determine C content. KMnO₄-C was measured as described by Blair et al. (1995), and the change in concentration of KMnO₄ was used to estimate the content of C oxidized, assuming that 1.0 mmol L⁻¹ of MnO₂⁻ was consumed (Mn⁷⁺ → Mn⁵⁺) in the oxidation of 0.75 mmol L⁻¹ (9.0 mg) of carbon.

The carbon management index (CMI) provides a reliable measure of the change in soil C dynamics of an experimental object relative to a reference object (Blair et al., 1995). CMI was determined as described by Blair et al. (1995). In our study, this index was calculated for the three afforestation areas using a reference sample value (the cropland) as follows:

\[
\text{CMI} = \frac{\text{Carbon pool index(CPI)} \times \text{Lability index(LI)}}{\text{Reference sample total organic C(g kg⁻¹)}} \times 100
\]

\[
\text{CPI} = \frac{\text{Experimental sample total organic C(g kg⁻¹)}}{\text{Reference sample total organic C(g kg⁻¹)}}
\]

\[
\text{LI} = \frac{\text{Lability of C(L)in experimental sample soil/Lability of C(L)in reference soil}}{1}
\]

\[
L = \text{KMnO}₄ - C/\text{(total organic C-KMnO₄-C)}
\]

### 3. Results

### 3.1. Total organic C concentrations and organic C stocks

Afforestation significantly affected total SOC concentrations and organic C stocks (Table 1). The total SOC concentrations were highest in the upper 0–5 cm soil layer and then decreased with increasing soil layer (Table 1). Afforestation with *R. pseudoacacia* and the mixed-species forest had significantly higher total organic C concentrations and stocks compared with cropland and *P. tabuliformis*. Furthermore, afforestation with *R. pseudoacacia*, and the mixed-species forest resulted in significant increases in total organic C concentrations in the 5–10 cm soil layer compared with cropland. Generally, no significant differences in total organic C concentrations were observed among the afforestation treatments with *R. pseudoacacia*, *P. tabuliformis*, and mixed stand in the 10–20 cm soil layer. Soil organic C stocks showed similar patterns as seen for total organic C concentrations.

### 3.2. Labile organic C fractions

Results for the 0–5 cm soil layer are described in the following paragraphs whereas results for the 5–10 cm and 10–20 cm layers are not repeated as they generally showed patterns similar to the 0–5 cm layer or no significant differences.

POC accounted for the largest proportion of labile C fractions (Fig. 1a). The three afforestation treatments significantly increased PO2 in the 0–5 cm soil layer compared with cropland. In comparison with the cropland alone, POC extracted from the *R. pseudoacacia*, *P. tabuliformis*, and the mixed stand was significantly increased by 57.4%, 22.2%, and 44.4% in the 0–5 cm, respectively. KMnO₄-C concentrations showed a pattern similar to POC among all land use types. KMnO₄-C in the 0–5 cm soil layer was highest for *R. pseudoacacia* (Fig. 1b). Afforestation with *R. pseudoacacia*, *P. tabuliformis*, and the mixed stand resulted in higher levels of KMnO₄-C compared with cropland, with concentrations in the 0–5 cm soil layer that were 2.3, 1.5, and 2.1 times higher than observed for cropland. MBC in the 0–5 cm layer was different among the three afforestation areas and cropland, with a significantly lower value observed for cropland than for the three afforestation areas (P < 0.01, Fig. 2a). MBC comprised the smallest proportion (1.6–3.1%) of labile C fractions. The three afforestation areas had significantly higher DOC compared with cropland (P < 0.01, Fig. 2b). The proportion of DOC varied from 3.1 to 5.2% of labile C fractions. The impacts of different afforestation compositions on the proportion of DOC were similar to the other organic C fractions described above, with the highest concentration observed for *R. pseudoacacia* and the lowest concentration observed for cropland.

### 3.3. Carbon management index

There were significant differences between the three afforestation areas and cropland in each soil layer for CMI, CPI, and LI (Table 2). Changes in CMI under the three afforestation treatments and cropland increased in the order of cropland < *P. tabuliformis* < mixed stand < *R. pseudoacacia*, with values ranging from 100 to 244 in the
4. Discussion

4.1. Effect of afforestation on soil total organic SOC

The distribution of total SOC concentration and stocks was affected by land use types (Zhao et al., 2014; Tong et al., 2016; Chen et al., 2018). In this study, total SOC concentration and stocks decreased with soil depth and were usually highest in the surface soil. Given that topsoil is the primary provider of soil nutrients and water for uptake by plant roots (Chen et al., 2016), this observation is not surprising. It was also found that afforestation led to a significant increase in total organic C, supporting our hypothesis. The average total SOC concentrations and stocks decreased in the order of $R$. pseudoacacia > $P$. tabuliformis > mixed stand > cropland. On the one hand, these differences could rationally be ascribed to diverse amounts of understory vegetation biomass returned to the soil (Wang et al., 2007; Kooch et al., 2019). For example, the $R$. pseudoacacia sites had the greater understory vegetation biomass amounts with the greater total soil organic C pool, while the $P$. tabuliformis sites had the lower understory vegetation biomass amounts, and consequently the lower total SOC (Fig. 3). On the other hand, deciduous broadleaf trees produce litter which is decomposed more rapidly than that of coniferous tree species (Wang et al., 2007; Sreekanth et al., 2013; Chen et al., 2018). Our study suggested that evergreen species utilize a more conservative carbon-use strategy than deciduous species. In addition, in line with our second hypothesis, soil C was significantly greater in the presence of $N_2$-fixing $R$. pseudoacacia than with $P$. tabuliformis. This was indicated by greater total SOC and labile organic C fractions in the soil surface layer (Table 1, Figs. 1, 2). It has been postulated that enhanced soil N availability in the $N_2$-fixing forest may increase total SOC and labile organic C fractions compared with the non-$N_2$-fixing forest (Butterbach-Bahl and Dannenmann, 2012).

The observation that both total SOC and labile organic C fractions in the 0–5 cm soil layer. Generally, afforestation with $R$. pseudoacacia resulted in the highest CMI value in each soil layer, indicating more labile C fraction.

![Fig. 1. Soil particulate organic carbon content (POC) (a) and potassium permanganate-oxidizable carbon (KMnO₄-C) (b) in the 0–5, 5–10, and 10–20 cm soil layers under different land-use types. Error bars represent standard deviations (n = 6). Different lowercase letters in the same soil layer indicate a significant difference at P < 0.05 among the different land-use types. CP: cropland; RF: Robinia pseudoacacia forest; PF: Pinus tabuliformis forest; RPF: Robinia pseudoacacia and Pinus tabuliformis forest.](image1)

![Fig. 2. Soil microbial biomass carbon content (MBC) (a) and dissolved organic carbon (DOC) (b) in the 0–5, 5–10, and 10–20 cm soil layers under different land-use types. Error bars represent standard deviations (n = 6). Different lowercase letters in the same soil layer indicate a significant difference at P < 0.05 among the different land-use types. CP: cropland; RF: Robinia pseudoacacia forest; PF: Pinus tabuliformis forest; RPF: Robinia pseudoacacia and Pinus tabuliformis forest.](image2)

![Table 2 Changes in carbon management index (CMI) at different soil layers under three afforestation treatments. Values represent mean ± SD for sample size of n = 6. Different lowercase letters in the same soil layer indicate a significant difference at P < 0.05 among the different land-use types. CF: cropland; RF: Robinia pseudoacacia forest; PF: Pinus tabuliformis forest; RPF: Robinia pseudoacacia and Pinus tabuliformis forest. CPI: carbon pool index; LI: lability index.](table)
the mixed stand were greater than in cropland but similar to the *R. pseudoacacia* stand indicated a non-synergistic effect of *R. pseudoacacia* intercropped with *P. tabuliformis* on SOC accumulation. Indeed, there is considerable potential for intercropping N\(_2\)-fixing species with woody non-N\(_2\)-fixing species to enhance SOC accumulation following afforestation (Yuan et al., 2016; Du et al., 2019). This may be explained by several factors, such as non-N\(_2\)-fixing species and soil texture (Wang and Xue, 1994). Therefore, the potential of intercropping with another locally and ecologically-important N\(_2\)-fixing species (*C. korshinskii*) with non-N\(_2\)-fixing species (*S. chelidophila* or *P. orientalis*) should be further studied.

### 4.2. Effect of afforestation on soil labile organic C fractions

Soil labile organic C fractions differed among different afforestation treatments and soil depths. Afforestation with different species compositions on the Loess Plateau positively increased POC, KMN\(_4\)-C, MBC, and DOC concentrations in the 20-cm soil profile, compared with cropland (Figs. 1, 2). It has been widely accepted that afforestation markedly increases soil labile organic C fractions (Deng et al., 2013; Sierra et al., 2013) directly or indirectly, which is consistent with our findings. There are several reasons for this observation. First, understory vegetation litter inputs may directly contribute to the soil labile organic C fractions (Fig. 3). Second, the positive microbial activities in the afforestation areas could increase the conversion of plant litter organic matter into labile forms of organic C (Poirier et al., 2013; Whalen et al., 2014). This result suggests that afforestation not only increased total SOC but also resulted in more soil labile organic C fractions (Table 1, Figs. 1, 2).

The CMI provides an indication of changes in the labile C dynamics of soil systems (Kalambukattu et al., 2013). Blair et al. (1995) reported that the actual CMI values make no sense, but relative differences in CMI values reflect how different management practices impact soil systems. In our study, CMI was highest under the three afforestation treatments and significantly larger than for cropland. *R. pseudoacacia* resulted in a higher CMI than *P. tabuliformis* and the mixed stand (Table 2). Our results agree with those from a long-term investigation by Blair et al. (2006). Tirol-Padre and Ladha (2004) explained that CMI variations observed under different soil management practices could be attributed to changes in organic matter quality, thus affecting the lability of C to KMnO\(_4\) oxidation. The KMnO\(_4\)-C concentration was also significantly affected by the forest expansion onto cropland, with the highest value for *R. pseudoacacia* and the lowest value for cropland (Fig. 1b). Similar concentrations of soil KMN\(_4\)-C have also been reported under long-term vegetation restoration (Zhao et al., 2014, 2015). Since CMI values indicate if an ecological system is in decline or being rehabilitated (Blair et al., 1995), the CMI values for cropland indicated deterioration, while afforestation fostered rehabilitation, especially for the *R. pseudoacacia* forest. Therefore, afforestation with *R. pseudoacacia* led to increases in KMnO\(_4\)-C as a percentage of total SOC and consequently increased CMI values compared with the agricultural management practices associated with crop production. These indices explain why afforestation not only increases soil total organic C, but also results in greater soil labile organic C fractions.

In addition, the POC level was substantially lower in cropland than in the three afforestation areas (Fig. 1a). POC contains decomposing plant litters with short-term turnover periods, which is consistent with the characteristics of total organic C concentrations (Feller and Beare, 1997). The three afforestation treatments resulted in the highest levels of POC because they produced the most plant residues. MBC and DOC are produced by the decomposition of soil organic matter mainly driven by soil microorganisms (Marschner and Bredow, 2002). Although they account for only a small proportion of total organic C (generally 1.6–3.1% for MBC and 3.1–5.2% for DOC) in the forest soils of our study, these two quantities are considered good indicators of the soil nutrient cycle (Moharana et al., 2012; Benbi et al., 2015). Significant increases in MBC and DOC were observed after afforestation, suggesting that forest expansion onto cropland had positive effects on the activity of microorganisms, probably by providing an available source of C substrate (Yang et al., 2012).

### 5. Conclusions

Our results clearly indicated that afforestation significantly increased total SOC concentrations, organic C stocks, labile organic C fractions (POC, KMN\(_4\)-C, MBC, DOC), and CMI compared with cropland in the Loess Plateau. The results also showed that afforestation with *R. pseudoacacia* resulted in greater carbon benefits than the other two afforestation treatments. Moreover, afforestation with *R. pseudoacacia* produced more total SOC and labile organic C fractions than either *P. tabuliformis*, the mixed stand, or cropland, mainly due to its ability to fix N\(_2\). Although afforestation increased total SOC, it also resulted in more soil labile organic C fractions in China’s Loess Plateau.

### CRediT authorship contribution statement

Qingyin Zhang: Conceptualization, Data curation, Software, Visualization, Writing - original draft. Xiaoxu Jia: Formal analysis, Funding acquisition, Project administration, Resources, Supervision, Validation. Xiaorong Wei: Data curation, Writing - review & editing. Mingan Shao: Funding acquisition. Tongchuan Li: Investigation, Methodology. Qiang Yu: Conceptualization, Writing - review & editing.

### Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

### Acknowledgments

This research was supported by the Fundamental Research Funds for the Central Universities (2452018088), the National Natural Science Foundation of China (41530854 and 41390461), the International Partnership Program of the Chinese Academy of Sciences (161461KYSB20170013). We thank the editors and reviewers for the constructive comments on the manuscript.

### Appendix A. Supplementary data

Supplementary data to this article can be found online at https://
References


Guo, L.B., Gi...